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1   **Previous land use and climate influence differences in soil organic carbon following**

2   **reforestation of agricultural land with mixed-species plantings**

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Abstract

Reforestation of agricultural land with mixed-species environmental plantings (native trees and shrubs) can contribute to mitigation of climate change through sequestration of carbon. Although soil carbon sequestration following reforestation has been investigated at site- and regional-scales, there are few studies across regions where the impact of a broad range of site conditions and management practices can be assessed. We collated new and existing data on soil organic carbon (SOC, 0-30 cm depth, $N = 117$ sites) and litter ($N = 106$ sites) under mixed-species plantings and an agricultural pair or baseline across southern and eastern Australia. Sites covered a range of previous land uses, initial SOC stocks, climatic conditions and management types. Differences in total SOC stocks following reforestation were significant at 52% of sites, with a mean rate of increase of 0.57 ± 0.06 Mg C ha^{-1} y^{-1}. Increases were largely in the particulate fraction, which increased significantly at 46% of sites compared with increases at 27% of sites for the humus fraction. Although relative increase was highest in the particulate fraction, the humus fraction was the largest proportion of total SOC and so absolute differences in both fractions were similar. Accumulation rates of carbon in litter were 0.39 ± 0.02 Mg C ha^{-1} y^{-1}, increasing the total (soil + litter) annual rate of carbon sequestration by 68%. Previously-cropped sites accumulated more SOC than previously-grazed sites. The explained variance differed widely among empirical models of differences in SOC stocks following reforestation according to SOC fraction and depth for previously-grazed ($R^2 = 0.18$-0.51) and previously-cropped ($R^2 = 0.14$-0.60) sites. For previously-grazed sites, differences in SOC following reforestation were negatively related to total SOC in the pasture. By comparison, for previously-cropped sites, differences in SOC were positively related to mean annual rainfall. This improved broad-scale understanding of the magnitude and predictors of changes in stocks of soil and litter C following reforestation is valuable for the development of policy on carbon markets and the establishment of future mixed-species environmental plantings.

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1. Introduction

There is increasing interest in reforestation of agricultural lands to sequester carbon in woody biomass and potentially mitigate greenhouse gas emissions (e.g. Canadell and Raupach, 2008; Cunningham *et al.*, 2015a). Reforestation can increase terrestrial carbon through humification and storage in soil organic carbon (SOC; Lal, 2005). While reforestation of agricultural lands significantly increases carbon sequestration in biomass compared with crop or pasture (e.g. Paul *et al.*, 2008; Cunningham *et al.*, 2015b), changes in SOC following reforestation are highly variable and uncertain, with increases, negligible change and decreases reported (e.g. Specht and West, 2003; Lima *et al.*, 2006; Harper *et al.*, 2012). Some variation may be explained by time since reforestation, as generally there are initial decreases in SOC stocks in the first five years after reforestation, followed by recovery to pre-establishment levels (approx. 10-30 years), and then a gradual increase (e.g. Paul *et al.*, 2002). However, most of the variability in SOC stocks following reforestation reflects differences in sequestration rates among climates, soil types, tree species and previous land uses (Guo and Gifford, 2002; Paul *et al.*, 2002; Laganière *et al.*, 2010).

Previous land use can be an important determinant of sequestration of SOC following reforestation, with increases in stocks on ex-cropland and predominantly losses on ex-pasture (Guo and Gifford, 2002; Laganière *et al.*, 2010). Climate can have a strong influence, with increases in tropical and sub-tropical regions compared with small decreases in temperate and Mediterranean-type regions (e.g. Paul *et al.*, 2002). Soils with high clay content generally having a larger capacity to accumulate SOC than those with lower clay content (Laganière *et al.*, 2010; Paul *et al.*, 2002). Further, the tree species planted can affect carbon sequestration, with increases in SOC stocks under some nitrogen-fixing acacia trees (Kasel *et al.*, 2011; Forrester *et al.*, 2013), although these effects may be species-specific (Hoogmoed *et al.*, 2014), whereas decreases in SOC stocks have been found under pines (Parfitt *et al.*, 1997; Turner and Lambert, 2000).

Accumulation of plant litter is an additional store of carbon in forests until a steady state between litterfall and decomposition is reached (Paul *et al.*, 2003). Rates of litter accumulation in

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81 native and plantation forests differ widely among forest types and species (e.g. Spain, 1984;

82 Adams and Attiwill, 1991; Fernandez-Nunez *et al.*, 2010), predominantly reflecting differences

83 in litter quality and climate (Prescott, 2010). Pine plantations can accumulate particularly thick

84 and recalcitrant litter layers (e.g. Paul *et al.*, 2003; Paul and Polglase, 2004) compared with those

85 under other plantation species (e.g. Turner, 1986; Harper *et al.*, 2012). Comparable studies under

86 mixed-species plantings are limited, but suggest that thick layers of up to approximately15 t dry

87 matter (DM) ha^{-1} can accumulate within two decades when eucalypts are planted (Cunningham *et*

88 *al.*, 2012).

89   SOC exists as a diverse mix of organic materials with different susceptibilities to biological

90 decomposition (Baldock *et al.*, 2013a). Reforestation may change the molecular form of SOC

91 and, consequently, increase the stability of the stock (Cunningham et al., 2015b). For a given

92 organic carbon content, the provision of energy to soil organisms should increase with increasing

93 proportion of plant litter-like components and decrease with increasing proportion of

94 biologically-recalcitrant charcoal or char-like components (Baldock *et al.*, 2013a). Reforestation

95 may increase inputs of more resistant SOC to soil but generally there is little increase in resistant

96 humic material within three decades, although earlier increases have been observed (e.g. Del

97 Galdo *et al.*, 2003; Cunningham *et al.*, 2015b). Understanding the form of SOC sequestered after

98 reforestation (i.e. its stability) is important in predicting the longer-term rates of sequestration and

99 resilience of carbon stocks to future change (e.g. with climate change), and to calibration and

100 verification of process-based models of turnover and accumulation of SOC.

101   Establishment of mixed-species environmental plantings (i.e. plantings of native tree and

102 shrub species established for environmental benefits with no intention to harvest) on agricultural

103 land can be an economically-viable option in lower rainfall (< 1000 mm y^{-1}) regions (Crossman

104 *et al.*, 2011; Polglase *et al.*, 2013). Indeed, environmental plantings are increasingly being

105 established to sequester carbon because of their co-benefits to the environment and biodiversity

106 (e.g. Mitchell *et al.*, 2012). Consequently, measurement and modelling of biomass carbon in

5

107   environmental plantings across a broad range of climatic and management conditions has been

108   the focus of recent work (e.g. Paul *et al.*, 2015). In contrast, there are limited measurements of

109   changes in soil and litter carbon under such plantings (e.g. Cunningham *et al.*, 2015b), and little

110   is known about their potential to sequester carbon in litter and soil compared with production

111   forests (Cunningham *et al.* 2015a). Global meta-analyses of soil carbon sequestration following

112   reforestation (Silver *et al.*, 2000; Guo and Gifford, 2002; Paul *et al.*, 2002; Laganière *et al.*, 2010)

113   include few studies of mixed-species plantings, and even meta-analyses of biomass accumulation

114   in mixed-species plantings have been dominated by plantings with only two species (Piotto,

115   2008; Hulvey *et al.*, 2013). Further, environmental plantings are highly variable, being

116   established across a much broader range of climates, previous land uses and landscape positions

117   than a given commercial plantation type (Paul *et al.*, 2015).

118   Here, we assessed potential predictors of soil carbon sequestration under environmental

119   plantings, which will inform their future establishment, calibration of carbon accounting models

120   and development of policy on carbon markets. A national dataset of 117 Australian sites was

121   collated and analysed, which represented much of the temperate and Mediterranean-type climates

122   across the continent. Three key research questions were addressed in relation to changes in

123   carbon sequestration following reforestation with environmental plantings:

124   1) Are there significant differences in stocks of total SOC and its fractions (particulate, humus,

125   resistant)?

126   2) Are estimates of carbon sequestration significantly increased when stocks of litter are

127   included?

128   3) What are the key site conditions and management practices that determine the magnitude of

129   differences in SOC and litter stocks?

130

2. Methods

2.1 Study sites

New and existing data were collated from 117 mixed-species environmental plantings (subsequently termed 'environmental plantings') established on agricultural land. Plantings were across southern and eastern Australia (latitude -30.9 to -38.7 °S, longitude 117.4 to 150.3 °E), and covered the range of rainfall zones where planting occurs (380 to 1147 mm y^{-1}; Table 1, Fig. 1). Thirty-six new sites were measured to improve the representativeness of plantings with respect to age, previous land use, productivity (aboveground biomass increment) and soil texture (Table 1). Environmental plantings are often established along stream banks to reduce erosion or in areas with shallow water tables to mitigate dryland salinity by minimising recharge (George *et al*., 1999). Thus sites were selected to include both dryland plantings across a range of landscape positions ($N = 97$) and riparian plantings ($N = 20$).

Previous land use at the sites included grazing, cropping, and rotational cropping and grazing (Table 1). For analysis, we combined cropping with rotational cropping and grazing as 'cropping'. There were fewer previously-cropped sites ($N = 32$) than previously-grazed sites ($N = 85$) reflecting current reforestation activity in Australia. Because environmental plantings are a relatively new land use, mean age of the plantings was 14 y, with 95% of plantings aged between 1 and 28 y. Biomass productivity varied greatly among sites (0.2 to 31.7 Mg DM ha^{-1} y^{-1}, calculated using methods of Paul *et al*., 2015) due to differences in site characteristics (e.g. soil type, rainfall), planting structure (e.g. stand density, species-mix, planting configuration) and age.

2.2 Soil and litter sampling at existing sites

Existing data on total organic carbon (TOC) in soil under environmental plantings were collated from 81 sites in temperate Australia. These data were collected between 2009 and 2012 using varying protocols including plot- and transect-based sampling. Studies limited to locations directly under trees were excluded because they are not representative of soil from across all

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157   planting strata (i.e. under trees, between trees along rows, inter-rows). Environmental plantings

158   were paired with the adjacent agricultural land use at 64 sites, whereas there had been a prior

159   baseline sampling (i.e. measurement before and after reforestation) at the remaining 17 sites.

160   Soil was sampled to 30 cm depth from either four (0-5, 5-10, 10-20, 20-30 cm) or two depths

161   (0-10 and 10-30 cm, or 0-5 and 5-30 cm depths) using a corer (internal diam. 42-73 mm). At each

162   site, 5-18 replicates (each comprising 1-6 cores) from 1-8 plots/transects were collected. Cores

163   were retained individually or bulked into 1-2 composite samples per transect, providing 5-18

164   samples per depth for the planting and 1-10 samples per depth for the adjacent agricultural land

165   for analysis.

166   Existing data on plant litter (≤ 25 mm diam.) were collated from 53 of these sites (Table 1). At

167   each site, 3-20 replicates from 1-36 quadrats (0.06-0.25 m^2) per plot/transect were sampled, with

168   quadrats processed either separately or bulked.

169

170   2.3 2.2 Soil and litter sampling at new sites

171   New data on SOC and plant litter under environmental plantings and agricultural pairs were

172   collected using standard protocols from an additional 36 sites and repeat sampling at the 17

173   existing baseline sites ($N = 53$ sites) in 2013. Sites were selected to fit the criteria that plantings

174   were: i) at least 5 years old, ii) established on previous agricultural land, and iii) paired

175   adequately with the adjacent agricultural land with respect to soil type, landscape position and

176   land-use history prior to planting.

177   The standard protocol (at 36 new 'standard' sites) involved sampling soil and litter from a

178   representative 0.4 ha area within the planting, and within the adjacent agricultural field. The 0.4

179   ha plot was divided into 40 sampling units of 10 m $\times$ 10 m, with one sampling position randomly

180   located per unit. For belts (10-40 m wide), sampling units were across the entire belt and for

181   blocks (> 40 m wide) a buffer of at least 5 m was left on each edge to minimise edge effects.

182   Both plots in the planting and agricultural pair were aligned along the same contour to minimise

183   differences in soil type and management history among the corresponding 40 sampling units in

184   the two land uses. Cores were sampled randomly within each unit to ensure that variation at sub-

185   plot scales due to past management practices (e.g. ripping and sowing lines, VandenBygaart,

186   2006) was captured in the estimate for the whole plot. Three additional sites were sampled more

187   intensively ('intensive' sites) as part of a study investigating the adequacy of sampling intensity

188   (Cunningham *et al.*, in review).

189   Soil was sampled using a corer (internal diam. 40-72 mm) to 30 cm depth with a minimum of

190   two depth intervals (0-10 and 10-30 cm), but with four depths (0-5, 5-10, 10-20, 20-30 cm) and

191   three depths (0-5, 5-10 and 10-30 cm) at some sites. Plant litter was sampled from a quadrat (0.1-

192   0.25 m^2) above each soil sampling location. At the standard sites, for each of the soil depths and

193   the litter layer, five randomly-selected cores/quadrats were bulked into one composite sample.

194   This provided eight composite samples for each soil depth for both the planting and for the

195   adjacent agricultural land. At the intensive sites, for each of the two soil depths and the litter

196   layer, 120 cores/quadrats in the planting and 56 in the adjacent agricultural land were analysed

197   individually for TOC/litter mass. Soil samples from the intensive sites were bulked randomly into

198   eight composite samples per depth (from 16 samples in the planting and seven in the adjacent

199   agricultural land) for analysis of the particulate (POC), humus (HOC) and resistant (ROC)

200   organic carbon fractions.

201

202   2.4 Sample processing

203   Soil samples were air-dried and sieved to ≤ 2 mm. High clay content soils were crushed

204   (predominantly using a jaw-crusher) prior to sieving to ensure that aggregates were not retained

205   in the > 2 mm fraction. Representative subsamples for analysis were obtained by careful mixing

206   either using a riffle box (Civilab, 12 × 13 mm slotted box) or by quartering. Air-dried litter

207   samples were separated into two components, fine (< 2 mm) and coarse (> 2-25 mm), by passing

208   through a 2 mm sieve. For 47 of the 53 existing sites with litter data, only a total litter mass was

    9

209   recorded. Subsamples were oven-dried at 60-70 $^{\circ}$C to constant mass. Total litter stocks were

210   calculated in Mg DM ha^{-1}.

211

212   2.5 Analysis of soil carbon content

213   TOC concentration was determined on a subsample of finely ground < 2-mm air-dried soil

214   using a CNS-2000 or CNS TruMac analyser (LECO Corporation, St. Joseph, MI, USA). For 35

215   of the existing sites (Cunningham *et al.*, 2015b), C concentrations of each sample were

216   determined from a subsample of 4–5 mg using catalytic combustion and thermal conductivity

217   (vario MICRO cube, Elementar Analysensysteme GmbH, Hanau, Germany). At a further 20

218   existing sites, a wet oxidation method (Heanes, 1984) was used to determine TOC concentration.

219   Where appropriate, samples were fizz-tested for the presence of carbonates with a few drops of 1

220   M HCl placed directly on the sample before analysis. Positive fizz samples were re-analysed with

221   the elemental analyser after pre-treatment with H_2SO_3 as a 5–6 wt% SO_2 solution to remove

222   inorganic carbon.

223   The POC, HOC and ROC fractions for all soil samples were predicted using mid-infrared

224   spectroscopy (MIRS) following the methods of Baldock *et al*. (2013a). The predictions were

225   made using new calibrations (Madhavan *et al.*, 2016) incorporating soils from environmental

226   plantings (largely this study) and agricultural soils (Baldock *et al.*, 2013a).

227   The moisture content in the < 2 mm soil samples after air drying was quantified

228   gravimetrically by oven drying a 30-40 g sample at 105 °C for 16-24 h. To estimate SOC

229   contents from soil carbon concentrations, volumetric soil mass (g cm^{-3}) of the 0-10, 10-30 and 0-

230   30 cm soil depths was calculated from the mass of the fine (< 2 mm) fraction and volume of the

231   bulked samples corrected for oven-dry (105 °C) moisture content. TOC, POC, HOC and ROC

232   stocks were calculated in Mg ha^{-1} for these fixed depths. Stocks of TOC were corrected for

233   equivalent soil mass following the equation given in Wendt and Hauser(2013). To do this, the

234   soil mass to 30 cm depth was calculated for all replicates and the 10th percentile of the masses

10

235   determined. The proportion of the soil mass required from the deepest depth interval (either 20–

236   30 cm or 10-30 cm) to achieve the target mass was applied to determine the mass of carbon

237   included from this depth increment.

238

239   2.6 Analysis of litter carbon content

240   A subset of 298 litter samples were selected for analysis of C content, including fine and

241   coarse fractions, from across 29 standard sites. Total C concentration was determined on a

242   subsample of finely-ground litter using a CNS-2000 or CNS TruMac analyser (LECO

243   Corporation, St. Joseph, MI, USA). Across the broader dataset ($N = 106$ sites), C stock of each

244   litter component was calculated in Mg C ha^{-1} either by multiplying the litter mass with the

245   measured C concentrations for the site ($N = 29$ sites), or for sites where litter C was not analysed,

246   using mean C concentrations ($N = 88$ sites).

247

248   2.7 Dataset collation

249   *2.7.1 Predictor variables*

250   To assess potential predictors of differences in SOC following reforestation, we collated a

251   range of site variables that were potential predictors. Collection of stand, management, climate

252   and soil variables are described below, with values of final candidate predictor variables used in

253   generalized linear models (see Section 2.8.2) given in Table 2.

254   Long-term (45-year; 1968-2012) monthly climate data for each of the sites were extracted

255   from the SILO Data Drill (Jeffrey *et al.*, 2001; https://www.longpaddock.qld.gov.au/silo/). We

256   derived both long-term climate variables and climate variables since planting (Table 2). Surrogate

257   data for climate, including elevation, latitude and longitude, were included as predictor variables.

258   Soil types and textures for each site were obtained using GIS layers of the Australian Soil

259   Classification system of the Atlas of Australian Soils (ABARES, 2004) and textural classes

260   (McKenzie *et al.*, 2000). At the broadest classification, the majority of the sampled soils (68%)

11

261 were Sodosols, so this was included as a predictor (Table 2). Soil texture was a categorical

262 variable with values of 1 (sand), 2 (sandy loam), 3 (loam), 4 (clay loam) or 5 (clay). Soil clay

263 content was obtained from spatial layers underpinning the national carbon accounting model

264 FullCAM (DIICCSRTE, 2014).

265   Land managers provided information about the age of the planting at the time of measurement

266 and the previous land use. Plantings were either planted as tubestock or direct seeded, and we

267 included a predictor for tubestock planting. Mean planting width was measured on site or

268 estimated from imagery in Google^{TM} Earth.

269   Measurement plots (N = 1-51 per planting, area = 0.01-3.37 ha) or transects (N = 2-8 per

270 planting, area = 0.001-0.010 ha) were established at each site to estimate biomass and stand

271 structure. Plots/transects were surveyed to coincide with soil sampling (2013 and 2009-2012,

272 respectively). Within each plot, stem diameters of all trees (at 130 cm height) and shrubs (at 10

273 cm height) were measured. Above-ground (AG) biomass for individual trees or shrubs was

274 estimated by applying existing species-specific allometric equations developed from

275 environmental plantings in temperate Australia (Paul *et al.*, 2013) and scaled to a live AG

276 biomass (Mg DM ha^{-1}) for each planting (Paul *et al*., 2015). Stand density was calculated as the

277 number of trees/shrubs per hectare. The proportion of trees, proportion of eucalypts (predominant

278 trees in plantings), proportion of acacias (dominant understorey in most plantings) and the

279 proportion of potential nitrogen-fixing trees (trees belonging to the genera *Acacia*, *Allocasuarina*

280 and *Casuarina*) were calculated based on stand density (Table 2). Proportions of tree types were

281 based on density as proportions based on live AG biomass were found to be poorer predictors

282 (data not shown).

283   Many riparian plantings, and some plantings from lower landscape positions, had

284 exceptionally high rates of accumulation of AG biomass compared with a recent national survey

285 (Paul *et al*., 2015), suggesting access to additional water. Consequently, whether a planting was

286   riparian, was established in an upper landscape position or was assumed to have access to

287   additional water were included as predictor variables (Table 2).

288

289   *2.7.2 Adequacy of sampling intensity*

290   SOC is often highly variable (coefficient of variation = 0.1–0.6, Wilson *et al.*, 1997; Paz-

291   Gonzalez *et al.*, 2000), so adequate sampling of different soil depths, cores within plots and plots

292   across a site, is required to provide accurate estimates (Cunningham *et al.*, 2012). Concurrent

293   work on adequate sampling intensity at the intensively-measured sites (Cunningham *et al.*, in

294   review) found that a minimum of 30 cores were required to have a 95% probability of estimating

295   SOC within 10% of the population mean. We assessed differences at all individual sites (N =

296   117) and then compared these differences with only sites with adequate sampling intensity (i.e. >

297   30 cores per site; N = 71 sites; see Section 2.8.1 below). However, only sites with adequate

298   sampling intensity were included in analyses of potential predictors of SOC following

299   reforestation (see Section 2.8.2 below). The dataset was further reduced by excluding one site

300   with unexplained very high carbon stocks (ca 170 t C ha^{-1} to 30 cm depth), resulting in N = 70, 32

301   and 38 sites for analysis of all sites, previously-cropped and previously-grazed sites, respectively.

302

303   2.8 Data analysis

304   *2.8.1Differences in SOC between land uses at individual sites*

305   For each site, we used a *t*-test to assess whether TOC, POC, HOC or ROC stocks differed

306   significantly between the paired mixed-species planting and agricultural land use by depth

307   interval (0-10, 10-30, 0-30 cm). Differences between land uses were used to infer changes in

308   SOC following reforestation, with the assumption that the agricultural land use was in steady

309   state (i.e. any differences in SOC were attributable to the plantings). The number of samples

310   available for statistical analysis depended on whether the site was: i) an existing site (N = 5-10

311   composite samples or individual cores), ii) a new standard site (N = 8 composite samples) or iii) a

13

312   new intensive site (N = 56-120 individual cores). The dominant component of the resistant

313   fraction (ROC) measured here is charcoal (Baldock *et al.*, 2013b). Because no difference in ROC

314   would be expected over the time since reforestation covered in this study (Skjemstad and

315   Spouncer, 2003; Skjemstad *et al.*, 2004), ROC was not included in subsequent analyses of the

316   key factors influencing differences in SOC.

317

318   *2.8.2 Key factors influencing differences in SOC*

319   We tested whether difference in SOC stock (calculated as SOC in the planting minus SOC in

320   the agricultural pair/baseline) across sites was related to known predictors of change following

321   reforestation (i.e. planting age, SOC stock in the agricultural pair/baseline, annual rainfall and

322   previous land use) using linear regression. The exception to this was previous land use, which

323   was categorical (either 'cropping' or 'pasture'), so *t*-tests were used. The only significant linear

324   regression was between HOC and age, albeit a poor relationship (R^2 = 0.05, F = 4.93, P = 0.03, N

325   = 70). Overall, difference in SOC differed significantly between previously-cropped and

326   previously-grazed land in the surface soil (df = 68, t = 2.09, P = 0.04). Further, because both

327   'reference' SOC stock (i.e. TOC in the agricultural pair or baseline) and MAR differed

328   significantly (degrees of freedom (df) = 68, t > 2.33, P < 0.05) between previously-cropped and

329   previously-grazed land uses, in subsequent analyses, the dataset was separated into previously

330   cropped (N = 32) and previously grazed (N = 38) sites.

331   Generalized linear modelling (GLM) with Bayesian model averaging (BMA) was performed

332   using the 'bic.glm' function in the 'BMA' package of R (Raftery *et al.*, 2008) to determine what

333   stand, management, climate and soil variables were strong predictors of difference in SOC (TOC,

334   POC, HOC) stocks. BMA involves fitting multiple models and finding predictors with the highest

335   posterior probabilities of inclusion in good models. Predictor variables with a probability of

336   inclusion [Pr(inc)] > 75% are generally considered strong candidates for inclusion as 'key

337   predictors' of the response variable (Thomson *et al*., 2007). Linear coefficients of correlation (*r*)

        14

338   were determined using the Pearson's correlation coefficient and highly-correlated variables ($r >$

339   0.7) were excluded from this modelling because they influence the results of BMA analyses

340   (Thomson *et al.*, 2007). Values of final candidate predictor variables used in generalized linear

341   models are given in Table 2.

342   For all models assessed, the number of predictor variables was restricted to four (where $N =$

343   32-38) or seven (where $N = 70$) based on empirical evidence that the risk of over-fitting

344   regression models is high if the number of predictors is more than approximately $n/10$ (Harrell *et*

345   *al.*, 1984). Predictors were selected by initially running the model with all variables, and then

346   rerunning the model with only the seven/four variables with the highest Pr(inc). Model averaging

347   further reduces the potential risk of over-fitting because the probability that multiple models

348   represent spurious response–predictor relationships is substantially lower than the probability that

349   any single model is over-fitted (Thomson *et al*. 2007).

350   Using SOC stocks based on equivalent soil mass was explored but gave similar results, albeit

351   with poorer models ($R^2 = 0.01$-0.40, $N = 32$-70) when compared with those using SOC stocks

352   based on fixed depths ($R^2 = 0.07$-0.64, $N = 32$-70). Consequently, all analyses presented here are

353   for SOC stocks based on fixed depths.

354

355   **3. Results**

356   3.1 Differences in SOC between land uses: individual site analysis

357   When considering the total dataset ($N = 117$ sites), there was no statistically significant (P

358   <0.05) difference in total SOC (TOC; data not shown) between land uses for the majority of sites

359   (60%). Based on the assumption that any differences between land uses were attributable to the

360   plantings (see Section 2.8), and when only sites with adequate replication ($N = 71$) were included,

361   there were significant increases in TOC stocks (48% of sites), few significant decreases (4% of

362   sites) or no significant difference (48% of sites; see Table A1). Significant differences in TOC

363   stocks were more common at 0-10 cm depth than at 10-30 cm depth (Table A1), and in

15

364   previously-cropped than previously-grazed sites (data not shown). Across all 71 sites, the mean

365   ($\pm$SE) rate of increase in TOC stocks was 0.57 ± 0.06 Mg ha^{-1} y^{-1} at 0-30 cm depth (Table 3).

366   POC increased significantly at 46% of sites, compared with 27% of sites for the humus (HOC)

367   fraction. HOC comprised the largest proportion of TOC (0.55 ± 0.01 and 0.53 ± 0.01), followed by

368   ROC (0.25 ± 0.01 and 0.25 ± 0.01) and POC (0.19 ± 0.01 and 0.22 ± 0.01), for the agricultural

369   pair/baseline and the plantings, respectively (data not shown). The largest relative difference was

370   observed for the labile fraction, with a mean rate of increase in POC stock of 3.1 ± 0.5 % y^{-1}

371   compared with 1.2 ± 0.3 % y^{-1} for HOC stocks (Table 3). Across the 71 sites, mean rates of

372   increase in stocks of POC and HOC were 0.19 ± 0.04 and 0.21 ± 0.06 Mg ha^{-1} y^{-1}, respectively

373   (Table 3).

374

375   3.2 Key factors influencing differences in SOC

376   *3.2.1 Effects of previous land use*

377   Mean annual difference in TOC stock was higher for former cropped than grazed land at 0-10

378   cm depth ($df = 68$, $t = 2.09$, $P = 0.04$; Fig. 2a) but not for 0-30 cm depth ($df = 68$, $t = 1.06$, $P =$

379   0.29). Conversely, TOC stocks in the agricultural pair or baseline were significantly lower for

380   previously-cropped than previously-grazed sites both at 0-10 cm ($df = 68$, $t = 2.48$, $P = 0.02$) and

381   0-30 cm depth ($df = 68$, $t = 2.33$, $P = 0.02$; Fig. 2b). However, unlike TOC, there was no

382   significant difference between previous land uses for mean annual difference in stocks of POC or

383   HOC ($df = 68$, $t < 0.80$, $P > 0.42$).

384

385   *3.2.2 Effects of multiple predictor variables across sites*

386   For the previously-grazed sites, 'reference' TOC (i.e. TOC in the agricultural pair or baseline;

387   Fig. 3a) and temperature were strong predictors of the difference in TOC at 0-10 cm depth, with

388   differences becoming less pronounced on sites with higher reference (or initial) TOC (Table 4).

389   When assessing the difference in TOC at 0-30 cm depth, 'reference' TOC remained a strong

390   predictor but the models were on average poor ($R^2 = 0.24$). For the previously-cropped sites, the

391   strong predictors of difference in TOC were long-term mean annual rainfall and its interaction

392   with planting age (Table 4). Differences in TOC increased with stand age, but with increases only

393   being pronounced at sites where MAR was > 550 mm y^{-1} (Fig. 3b). The model for difference in

394   TOC at 0-10 cm depth was similar to that at 0-30 cm, but was poor ($R^2 = 0.14$) at 10-30 cm depth

395   with no strong predictors (Table 4).

396   Because the largest differences in SOC fractions were observed in the surface soil, only

397   models for 0-10 cm are presented for differences in POC and HOC (Table 5). Across the

398   previously-cropped sites, strong predictors of difference in POC and HOC in the surface soil

399   included long-term MAR, minimum temperature and measures of stand composition including

400   the proportions of acacia trees and nitrogen-fixing trees (Table 5). Long-term MAR, minimum

401   temperature and the proportion of acacia trees had a positive influence, while the proportion of

402   nitrogen fixing trees had a negative effect. By comparison, across the previously-grazed sites,

403   'reference' SOC was the only strong predictor of difference in POC while 'reference' SOC and

404   climatic factors including long-term MAR and long-term maximum temperature were strong

405   predictors of difference in HOC (Table 5). 'Reference' SOC had a negative influence on

406   difference while long-term MAR and long-term maximum temperature had a positive effect.

407

408   3.3 Change in litter following reforestation

409   Across all 106 sites with litter data (53 existing, 53 new), total litter stocks in the plantings

410   ranged from 2.4 to 30.6 Mg DM ha^{-1}, with rates of litter accumulation ranging from 0.22 to 2.27

411   Mg DM $ha^{-1}\,y^{-1}$. Similarly, across the 65 sites corresponding to sites with adequate sampling

412   intensity for SOC, mean accumulation rates of litter dry matter and litter carbon were 0.94±0.06

413   Mg DM $ha^{-1}\,y^{-1}$ and 0.39±0.02 Mg C $ha^{-1}\,y^{-1}$, respectively. When the combined change in soil and

414   litter stock following reforestation was considered at the 65 sites with adequate sampling and

17

415   litter data, mean rates of accumulation were 1.25 ± 0.20 and 0.79 ± 0.12 Mg C ha^{-1} y^{-1} for

416   previously-cropped ($N = 30$) and previously-grazed sites ($N = 34$), respectively (Fig. 2c).

417

418   **4. Discussion**

419   4.1 Change in SOC stocks following reforestation

420   Reforestation of agricultural land with environmental plantings resulted in a difference in

421   stocks of total SOC (0-30 cm) of 0.57 ± 0.06 Mg C ha^{-1} y^{-1} (Table 3). Significant differences in

422   total SOC (sequestration) following reforestation were found at 52% of sites; for the majority of

423   these total SOC increased, while total SOC decreased at 4% of sites (Table A1). Differences in

424   total SOC were largely a result of differences in the POC fraction with significant increases at

425   48% of sites, compared only 27% of sites for the HOC fraction (Table A1). Relative increase was

426   highest for POC but because HOC comprised the largest proportion of total SOC of all fractions,

427   absolute differences in POC and HOC were similar (Table 3).

428   The increase in POC found here agrees with studies of afforestation of grassland across a

429   rainfall gradient in South America (Berthrong *et al.*, 2012), and of cropland in Midwestern USA

430   (DeGryze *et al.*, 2004). There are few comparable studies of HOC, as most report on 'recalcitrant

431   C', which is equivalent to the combined humus and resistant fractions in our study. No change in

432   recalcitrant C following reforestation was found in young deciduous *Populus* (DeGryze *et al.*,

433   2004) or evergreen *Eucalyptus-Acacia* plantings (Kasel *et al*., 2011). However, increases in

434   stocks of recalcitrant SOC have been reported (Del Galdo *et al.*, 2003; Cunningham *et al.*,

435   2015b). Carbon in the humus fraction is considered to be more stabilised against microbial

436   decomposition due to formation of organo-mineral complexes (Baldock and Skjemstad, 2000).

437   By comparison, the resistant fraction is highly-carbonised organic material such as charcoal

438   (Baldock and Skjemstad, 2000) which can survive for > 500 years (Lehmann *et al.*, 2008) and

439   was therefore assumed not to change over the timeframe considered here.

18

440   There is little evidence for substantial change in soil carbon in the first three decades after

441   reforestation of former pastures (e.g. Post and Kwon, 2000; Guo and Gifford, 2002; Paul *et al.*,

442   2002; Hoogmoed *et al.*, 2012; Prior *et al.*, 2015). Our survey showed no effect of planting age on

443   SOC stocks, although differences ranged from increases, to negligible differences, to decreases at

444   a small number ($n = 4$) of sites (Table A1). Similarly, there were no consistent changes in soil

445   carbon following reforestation with environmental plantings aged up to 45 years (Cunningham *et

446   al*., 2015b). Further, a review of reforestation studies, largely of *Eucalyptus* species in Australia,

447   found little evidence for changes in soil carbon the first three decades (Hoogmoed *et al.*, 2012),

448   but increases in SOC were observed in some younger (<10-year-old) plantings. These findings

449   suggest that there may be site-specific factors influencing change in SOC following reforestation

450   with environmental plantings. Further, which factors are important is likely to change depending

451   on the scale (i.e. within regions e.g. Cunningham *et al*., 2015b, nationally e.g. this study, or

452   globally e.g. Paul *et al*., 2002) at which changes in SOC are considered.

453   We believe the findings of our survey are robust because of adequate sampling and pairing of

454   the plantings and adjacent fields. It is likely that many previous studies were not adequate in

455   terms of sampling intensity (see Cunningham *et al.*, in review), so they were more likely to be

456   inconclusive. We attributed any differences in SOC between the agricultural land use and the

457   plantings at the paired sites to be due to reforestation based on the assumption that the

458   agricultural land use is in steady state, i.e. management practices in the agricultural pair did not

459   change after reforestation. We believe this assumption to be reasonable and, at least in part,

460   verified by management histories in the 10 years prior to sampling collected for each of the

461   agricultural pairs. Further, similar significant increases in TOC were found for 13 of the 17 sites

462   with baseline and repeat sampling. With respect to pairing, recent detailed comparison of the

463   mineralogy using the MIR spectra of soils from the individual paired sites used here has shown

464   that the pairs share the same mineralogy, supporting the assumption that the sites were paired

465   adequately with respect to soil type (Madhavan, *unpublished*). This, combined with the filtering

19

466   of the dataset to include only sites with adequate sampling intensity, gives support to the

467   robustness of the observed changes in SOC.

468

469   4.2 Predictors of change in SOC following reforestation

470   *4.2.1 Effects of previous land use*

471   Previous land use had a significant effect on changes in soil carbon following reforestation

472   (Fig. 2). SOC stocks are generally higher in grazed than cropped land uses at a specific location

473   (e.g. Davy and Koen, 2013) and across regions (Rabbi *et al.*, 2014; Hobley *et al.*, 2015). At a

474   specific location, this difference is largely related to the higher disturbance (from tillage) of soil

475   under cropping compared with pasture soils, which have higher retention of carbon in aggregates

476   and more inputs of plant residues from roots (Percival *et al.*, 2000). Across regions, cropped

477   lands tend to be in areas of lower rainfall and, within these areas, we might expect increased SOC

478   with increased rainfall, particularly when increased rainfall coincides with cooler temperatures.

479   Here, the significantly lower SOC stocks in soils under cropping than grazing (Fig. 2b) was

480   consistent with earlier surveys (e.g. Paul *et al*., 2002). Because cropped lands are most depleted

481   in SOC in the surface soil, reforestation of previously-cropped land is likely to lead to larger

482   increases at the surface. In contrast, the surface (and total profile) of grazed lands is less likely to

483   be depleted (e.g. see Hobley *et al*., 2015), so the likelihood of soil carbon sequestration at these

484   sites is lower. Our survey showed that the magnitude of change in SOC following reforestation

485   was significantly higher for previously-cropped than previously-grazed sites in the 0-10 cm soil

486   depth (Fig 2a). This agrees the findings of with existing meta-analyses (Laganiere *et al*., 2010;

487   Guo and Gifford 2002; Paul *et al*. 2002), and suggests that sites with lower initial carbon (due to

488   greater depletion) have a greater capacity to sequester carbon following reforestation.

20

490   *4.2.2 Effects of multiple predictor variables across sites*

491   The largest proportional increases in SOC following reforestation are generally found in lower

492   rainfall areas (< 1200 mm y^{-1}; Guo and Gifford, 2002), such as those in this survey. Long-term

493   mean annual rainfall was a strong predictor of difference in TOC for 0-10 and 0-30 cm depths for

494   previously-cropped (but not previously-grazed) sites (Table 4), with a trend of increasing

495   difference in TOC stocks with increasing rainfall for the 0-30 cm layer (Fig. 3b). Similarly, for

496   previously-cropped (but not previously-grazed) sites, difference in POC was positively related to

497   MAR in the 0-10 cm layer (Table 5). A survey of agricultural lands in eastern Australia also

498   found that SOC in surface soil was driven predominantly by climate, particularly precipitation,

499   vapour pressure deficit and relative humidity (Hobley *et al.*, 2015). Despite the rainfall effect at

500   previously-cropped sites, other variables potentially affecting water availability, such as access to

501   groundwater or landscape position, generally were not strong predictors of change in SOC stocks

502   (Tables 4 and 5).

503   Difference in TOC in the surface soil was related negatively to SOC in the agricultural soil on

504   previously-grazed sites, but not for previously-cropped sites (Table 4; Fig 3a). This suggests that

505   reforestation of more productive pasture may lead to a loss in SOC. More productive pastures are

506   likely to have higher initial SOC stocks, which have the potential to increase losses in SOC

507   stocks observed in early decades after reforestation (Paul *et al*., 2002). Soils with high carbon

508   stocks may have higher nitrogen stocks and differing microbial communities (e.g. Fierer *et al.*,

509   2009), which play a key role in organic matter decomposition. Reports of the effects of

510   reforestation by environmental plantings on soil microbial communities are limited, but a recent

511   study of plantings on former pastures in relation to nitrogen supply found that differences in

512   microbial communities were specific to the tree species planted (Hoogmoed *et al*., 2014).

513   The species of trees planted are important to soil carbon sequestration following reforestation,

514   with the largest increases under broadleaf species (27%), intermediate responses under eucalypts

      21

515 and pines (12%) and little change under other conifers (2%, Laganière *et al.*, 2010). Plantings that

516 include nitrogen-fixing tree species can have higher productivity (Forrester *et al.*, 2006) and

517 higher retention of original soil carbon stocks (Resh *et al.*, 2002), suggesting their inclusion may

518 accelerate accumulation of SOC (Cunningham *et al.*, 2015a). Here, there were some species

519 effects on differences in HOC across the previously-cropped sites, with significant predictors

520 including the proportion of acacia trees and proportion of N-fixing trees (Table 5). The

521 proportion of acacia trees had a positive influence on change in HOC fractions in the surface 10

522 cm, suggesting that inclusion of these in plantings may increase sequestration of more stable

523 forms of carbon on previously-cropped sites. Counter-intuitively, differences in these fractions

524 were negatively related to the proportion of nitrogen fixing trees in the planting (Table 5).

525 Similarly, previous studies in environmental plantings comparing changes in SOC between fixers

526 and non-fixer trees found that effects were restricted to specific *Acacia* species (Hoogmoed *et al.*,

527 2014; Kasel *et al.*, 2011).

528

529 4.3 Changes in litter stocks following reforestation

530   Total litter stocks under young ($\leq$ 29-year-old) environmental plantings ranged from 2.4 to

531 30.6 Mg DM ha^{-1}. This is comparable with reported litter stocks, which range from 4.1-24.6 Mg

532 DM ha^{-1} in eucalypt-dominated open forests and woodlands (Pressland, 1982; Hutson, 1985;

533 Adams and Attiwill, 1991; Crockford and Richardson, 1998) and 19-34 Mg DM ha^{-1} under 26-

534 year-old farm forestry plantations of eucalypts (Harper *et al.*, 2012) across similar rainfalls.

535 Accumulation of litter C varied widely among sites, with the mean ($\pm$SE) rate of accumulation of

536 0.39$\pm$0.02 Mg C ha^{-1} y^{-1}, which was similar to the rate in soil (0.57$\pm$0.06 Mg C ha^{-1} y^{-1}). When the

537 additional changes in litter C were added to the changes in soil carbon, rates of C sequestration

538 were in the order of 1.25$\pm$0.20 and 0.79$\pm$0.12 Mg C ha^{-1} y^{-1} for previously-cropped and

539 previously-grazed sites, respectively (Fig. 2d).

540

22

541 4.4 Conclusion

542 There is potential to increase SOC stocks following reforestation with mixed-species

543 environmental plantings in the first two decades. Reforestation on former cropping sites has a

544 higher capacity to sequester SOC, at least in the surface soil, than on former grazing sites.

545 Differences in SOC stocks were generally predicted weakly by empirical models, suggesting that

546 either the factors driving these changes may be largely site-specific, or that predictors not

547 included here are more important. For example, there is potential for other processes, such as

548 increased deposition and reduced erosion, to increase site-based SOC in these agricultural

549 landscapes (Bird *et al.*, 1992).

550 Factors that were correlated strongly with difference in SOC stocks were predominantly long-

551 term rainfall at previously-cropped sites, and SOC stocks in the agricultural pair/baseline and

552 temperature at previously-grazed sites. This improved understanding of the potential predictors of

553 sequestration of soil carbon following reforestation across a broad range of site conditions and

554 management practices suggests that rates could be improved by planting on previously-cropped

555 sites in areas where rainfall is > 550 mm yr^{-1}. However, this does not account for the rates of

556 carbon sequestration in above-ground biomass and debris, and the current focus of reforestation

557 on previously-grazed sites, presumably due to demand for prime cropping land. More targeted

558 sampling to test specific hypotheses and drivers, such as addressing differences between former

559 cropped and grazed land uses, particularly in older plantings (> 30 years old) is the important

560 next step in building accurate models of carbon sequestration in soil following

561 reforestation. Experimental plantings are required to investigate the effects of species mix and

562 soil type on soil organic carbon following reforestation with mixed-species environmental

563 plantings.

564

23

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Table 1 Summary of the coverage and source of the datasets collated, including sampling method (paired = paired with agricultural land use, repeat = repeat sampling from agricultural baseline), age, previous land use (G = grazing, C = cropping, including rotational cropping and grazing), productivity of the plantings (above-ground biomass (AGB) increment; low = < 4 Mg DM ha^{-1} y^{-1}, high = ≥ 4 Mg DM ha^{-1} y^{-1})[1] and soil texture (light = ≤ 20% clay, heavy = > 20% clay)[2].

Type	SOC (*n* sites)	Litter (*n* sites)	Sampling	Age (y)	Previous land use (*n* sites)		AGB increment (*n* sites)		Soil texture (*n* sites)		Source
					G	C	Low	High	Light	Heavy	
Existing	20	11	paired	1-19	20	-	11	9	5	15	Read (2015)
Existing	35	35	paired	5-45	35	-	20	15	30	5	Cunningham *et al*. (2015b)
Existing	2	-	paired	15-20	-	2	2	-	2	-	T. Herrmann, *pers. comm.*
Existing	7	7	paired	8-16	5	2	2	5	7	-	T. Baker, *pers. comm.*
Existing	17	-	baseline	8-20	3	14	14	3	-	17	B. Wilson, *pers. comm.*
New	53	53	paired, repeat[5]	6-29	28	28	33	23	15	41	This study
Total	**117**[6]	**106**		**1-45**	**85**	**32**	**66**	**51**	**58**	**59**	

[1] Based on Paul *et al*., (2015);

[2] Based on spatial layers underpinning the National Carbon Accounting Toolbox (DIICCSRTE 2014);

[3] Repeat sampling of 17 baseline sites (B. Wilson);

[4] Totals exclude baseline sampling.

792 **Table 2** Variables used as candidate predictors of difference in soil organic carbon stocks and

793 litter stocks following reforestation with mixed-species plantings ($N = 117$ sites).

Code	Mean	SD	Min	Max	Definition
Climate & water availability					
MAR[a]	647	135	380	1147	Long-term mean annual rainfall (mm)
PLEVAP[a]	1467	228	1063	2681	Mean annual evaporation since planting (mm)
MAXT[a]	21.1	1.9	17.0	25.6	Long-term mean monthly max. temp. (°C)
MINT[a]	8.5	1.3	5.9	11.5	Long-term mean monthly min. temp. (°C)
PLMAXTHOT[a]	31.2	2.3	23.9	35.3	Mean max. temp. of the hottest month since planting (°C)
PLMINTCOLD[a]	2.6	1.3	-0.3	6.0	Mean min. temp. of the coldest month since planting (°C)
ELEV[a]	372	209	41	1055	Elevation (m)
RIPARIAN[b]	-	-	-	-	Riparian landscape position
UPPER[b]	-	-	-	-	Upper landscape position
ADDWATER[b]	-	-	-	-	Access to additional water
Soils					
AGSOILTC0-30[c]	46.5	20.2	15.4	160.2	Total soil carbon content of agricultural pair, 0-30 cm depth (Mg ha^{-1})
AGSOILTC0-10[c]	25.7	10.9	8.9	90.4	Total soil carbon content of agricultural pair, 0-10 cm depth (Mg ha^{-1})
AGSOILTC10-30[c]	21.3	11.2	6.5	69.8	Total soil carbon content of agricultural pair, 10-30 cm depth (Mg ha^{-1})
CLAY[d]	26.6	10.0	5.0	53.0	Soil clay content (%)
TEXTURE[e]	2.9	1.0	1.0	5.0	Soil texture (1-5)
SODOSOL[b,e]	-	-	-	-	Sodosol soil type
Previous land use and age					
GRAZED[b]	-	-	-	-	Previous land use grazing
CROPPED[b]	-	-	-	-	Previous land use cropping
AGE	14	7	1	45	Planting age at sampling (y)
Management					
TS[b]	-	-	-	-	Planting established using tubestock
WIDTH	64	63	7	416	Mean planting width (m)
DENSITY[c]	1951	3269	124	20768	Density of trees and shrubs (ha^{-1})
PROPEUC[c]	0.53	0.25	0.00	1.00	Proportion of eucalypts
PROPTREE[c]	0.75	0.20	0.10	1.00	Proportion of trees
PROPACA[c]	0.31	0.25	0.00	1.00	Proportion of acacias (trees and shrubs)
PROPACATREE[c]	0.16	0.18	0.00	0.83	Proportion of acacia trees
PROPFIXER[c]	0.21	0.18	0.00	0.83	Proportion of N fixing trees
LIVEAGB[c]	62.8	57.3	0.4	411.5	Live aboveground biomass (Mg DM ha^{-1})
LAT[f]	-35.45	1.78	-38.67	-30.86	Longitude (° E)
LONG[f]	146.1	4.5	117.4	150.3	Latitude (° S)

794 [a] Derived from SILO data drill 1968-2012 (Jeffrey et al. 2001; https://www.longpaddock.qld.gov.au/silo/); [b] Binary categorical

795 variable with value of 0 or 1; [c] Derived from plot measurements in this study; [d] Derived from FullCAM (DIICCSRTE 2014); [e]

796 Derived from GIS layers in the Australian Soil Classification system of the Atlas of Australian Soils (ABARES 2004) and textural

797 classes (McKenzie et al. 2000); [f] Geographic coordinates not used in main models.

798 **Table 3** Difference in soil organic carbon (ΔSOC) following reforestation with mixed-species

799 plantings at sites with adequate sampling intensity ($N = 71$, plantings aged 2-29 y). ΔSOC values

800 are mean annual percentage change (% y^{-1}) and mean annual change (Mg C ha^{-1} y^{-1}); standard

801 errors of the means are given in parentheses.

SOC fraction	Depth (cm)	ΔSOC	
		% y^{-1}	Mg ha^{-1} y^{-1}
Total	0-30	1.5 (0.3)	0.57 (0.10)
(TOC)	0-10	1.4 (0.3)	0.21 (0.06)
	10-30	1.8 (0.5)	0.35 (0.10)
Particulate	0-30	3.1 (0.5)	0.19 (0.04)
(POC)	0-10	3.0 (0.5)	0.09 (0.02)
	10-30	4.9 (1.4)	0.09 (0.02)
Humus	0-30	1.2 (0.3)	0.21 (0.06)
(HOC)	0-10	1.1 (0.3)	0.09 (0.03)
	10-30	1.5 (0.5)	0.12 (0.04)

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804 **Table 4** Results of generalized linear modelling with Bayesian model averaging of difference in

805 total organic carbon (ΔTOC) for fixed depths (0-10, 10-30, 0-30 cm) in relation to stand,

806 management, climate and soil variables at sites with adequate sampling intensity, by previous

807 land use. Only strong predictors Pr(inc) > 0.75 are shown, with the exception of predictors

808 included with interactions (italicized). Predictor variable codes are defined in Table 2.

Response variable	n	Mean model R^2	Predictor variable	Pr(inc)	Std. coefficient $\pm$ SD
Grazed					
ΔTOC, 0-30 cm	38	0.24	AGSOILTC0-30	0.76	-36.9 $\pm$ 41.3
ΔTOC, 0-10 cm	38	0.46	AGSOILTC0-10	1.00	-15.3 $\pm$ 3.8
			MINT	0.79	10.8 $\pm$ 7.4
ΔTOC, 10-30 cm	38	0.20	-	-	-
Cropped					
ΔTOC, 0-30 cm	32	0.52	MAR	1.00	100.1 $\pm$ 41.7
			MAR$\times$AGE	0.95	-94.4 $\pm$ 60.8
			AGE	*0.71*	*49.6 $\pm$ 47.8*
ΔTOC, 0-10 cm	32	0.53	MAR	1.00	72.3 $\pm$ 30.6
			MAR$\times$AGE	0.94	-65.0 $\pm$ 44.6
			AGE	*0.68*	*33.5 $\pm$ 35.0*
ΔTOC, 10-30 cm	32	0.14	-	-	-

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811 **Table 5** Results of generalized linear modelling using Bayesian model averaging of difference in

812 particulate (ΔPOC) and humus (ΔHOC) soil organic carbon fractions in the surface soil (0-10

813 cm) in relation to stand, management, climate and soil variables at sites with adequate sampling

814 intensity by previous land use. Only strong predictors Pr(inc) > 75 shown. Predictor variables are

815 defined in Table 2.

Response variable	n	Mean model r^2	Predictor variable	Pr(inc)	Std. coefficient ± SD
Previously-grazed					
ΔPOC	38	0.19	AGSOILTC0-10	0.90	-3.95 ± 2.01
ΔHOC	38	0.49	MAXT	1.00	25.56 ± 8.92
			MAR	1.00	10.46 ± 3.46
			AGSOILTC0-10	0.93	-5.92 ± 2.76
Previously-cropped					
ΔPOC	32	0.39	MAR	1.00	7.22 ± 2.06
ΔHOC	32	0.64	PROPACATREE	1.00	7.28 ± 2.23
			PROPNFIXTREE	1.00	-13.22 ± 2.69
			MINT	1.00	22.01 ± 6.95
			MAR	0.80	8.20 ± 5.49

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Appendix 1

818 **Table** Differences in soil organic carbon (ΔSOC) following reforestation with mixed-species plantings at sites with

819 adequate sampling intensity ($N = 71$). ΔSOC values are presented as mean annual percentage change (% y^{-1}) and

820 mean annual change (Mg C ha^{-1} y^{-1}) for significant increases (+), significant decreases (-) and no significant change

821 (0). Standard errors of the means are given in parentheses. POC = particulate organic carbon, HOC = humus organic

822 carbon.

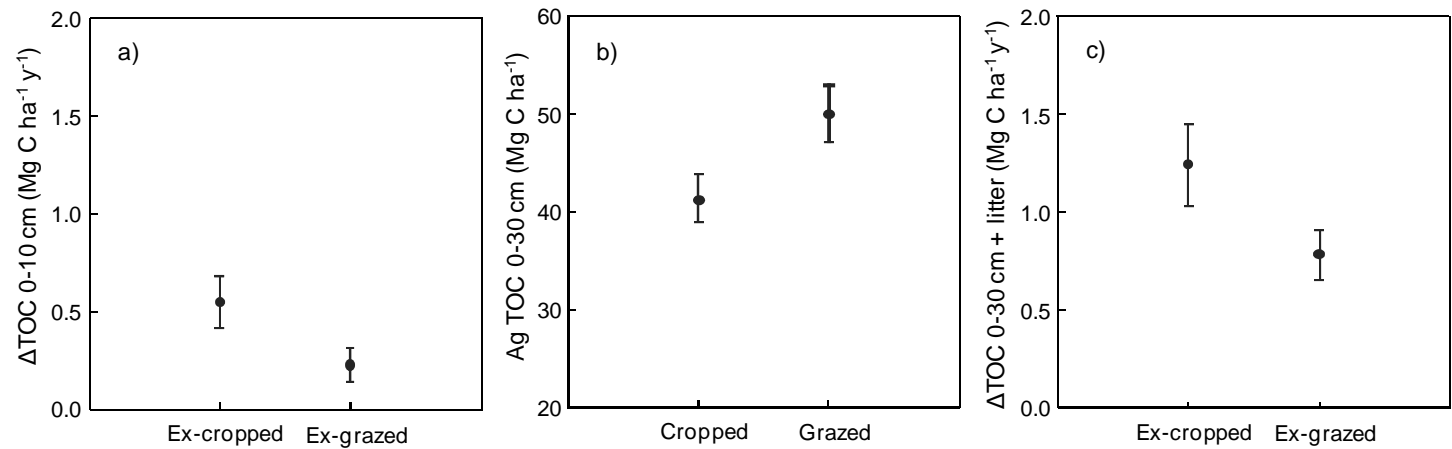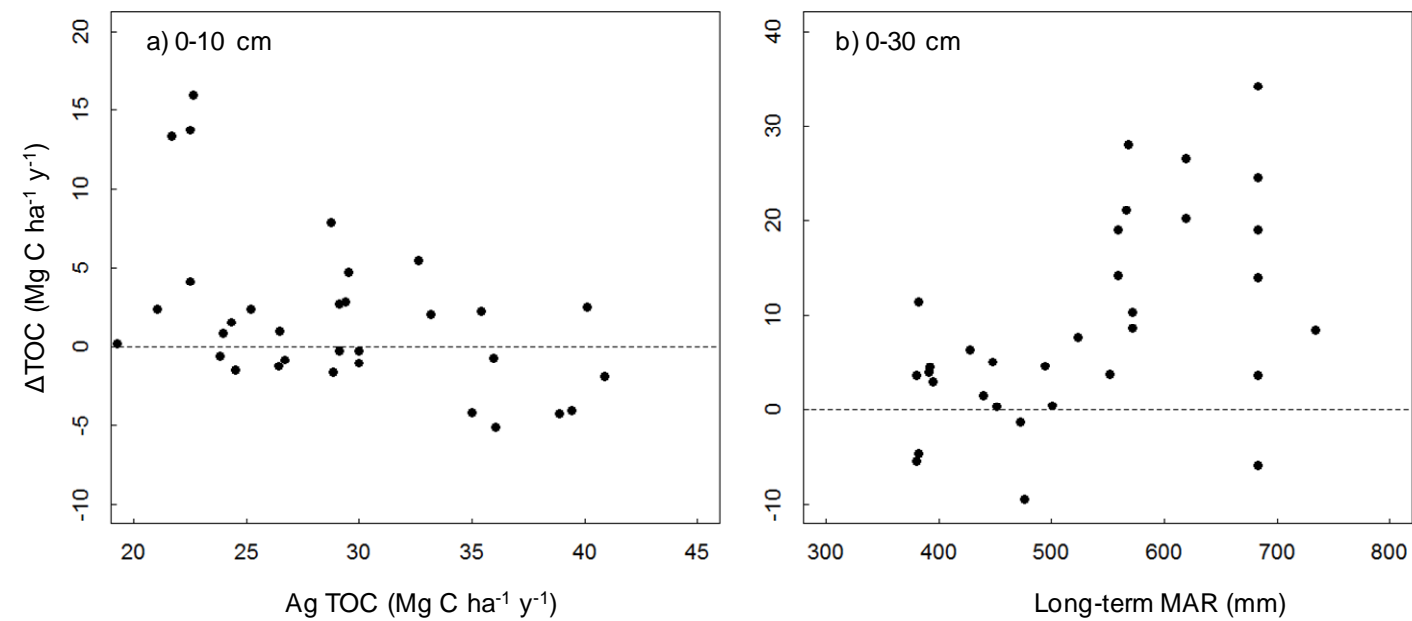| SOC fraction	Depth (cm)	Change	ΔSOC % y^{-1}	ΔSOC Mg ha^{-1} y^{-1}	*n*	Age range y
Total	0-30	+	2.7 (0.4)	1.1 (0.2)	34	2-23
		0	-	-	34	6-29
		-	-1.3 (0.2)	-0.9 (0.2)	3	9-18
	0-10	+	3.5 (0.5)	0.6 (0.1)	29	6-29
		0	-	-	36	2-21
		-	-1.4 (0.4)	-0.7 (0.4)	6	9-13
	10-30	+	4.3 (1.2)	0.9 (0.2)	25	2-23
		0	-	-	44	3-21
		-	-2.9 (1.2)	-0.3 (0.2)	2	18-29
POC	0-30	+	5.6 (0.8)	0.4 (0.1)	33	2-29
		0	-	-	33	3-21
		-	-1.8 (0.4)	-0.2 (0.1)	5	9-16
	0-10	+	5.5 (0.8)	0.2 (0.0)	28	8-29
		0	-	-	37	2-21
		-	-1.5 (0.2)	-0.2 (0.1)	6	9-16
	10-30	+	11.5 (2.0)	0.3 (0.1)	17	3-23
		0	-	-	46	2-21
		-	-5.1 (1.2)	-0.1 (0.0)	8	9-29
HOC	0-30	+	2.5 (0.3)	0.5 (0.1)	19	8-23
		0	-	-	47	2-29
		-	-1.3 (0.1)	-0.5 (0.1)	5	9-18
	0-10	+	4.0 (0.6)	0.3 (0.0)	20	8-29
		0	-	-	41	2-21
		-	-1.2 (0.2)	-0.2 (0.0)	10	9-18
	10-30	+	3.1 (0.4)	0.4 (0.1)	15	6-23
		0	-	-	48	2-21
		-	-1.8 (0.2)	-0.2 (0.0)	10	9-29

823

824   Figure 1
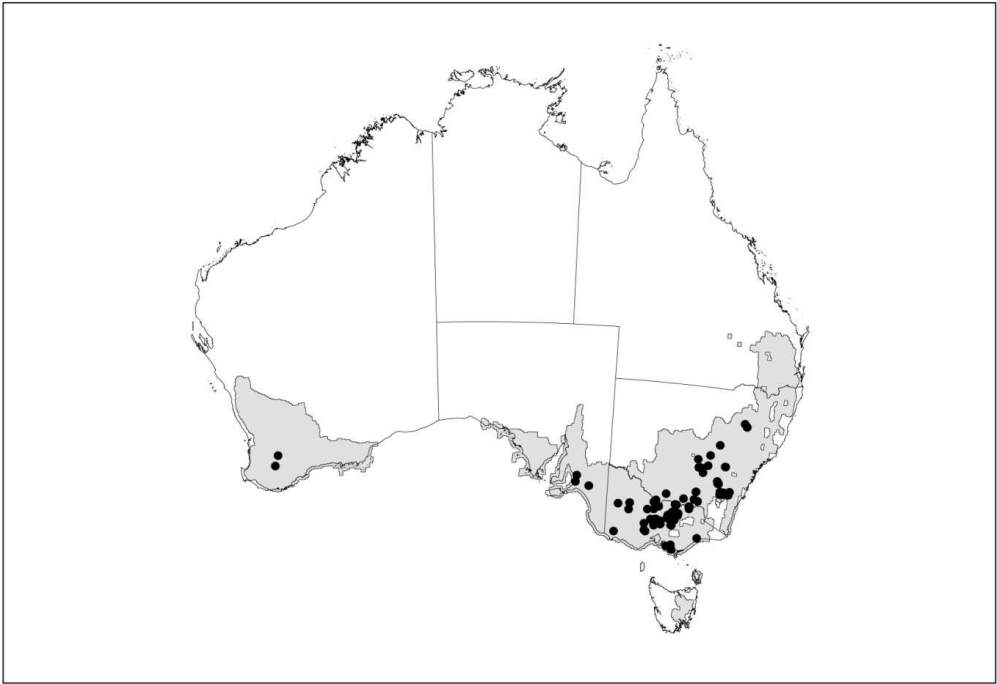
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827   Figure 2

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